

## Global surface ocean alkalinity climatology

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In many parts of the ocean alkalinity ( $A_T$ ) data are severely limited compared to the sea surface salinity (SSS) and temperature (SST) data sets, which are several orders of magnitude larger. Therefore, determining the global distribution of surface  $A_T$  is problematic. Empirical algorithms that relate surface  $A_T$  to SSS and SST are particularly useful in constructing the global distribution of  $A_T$  when combined with the global fields of SSS and SST.

The first set of global relationships of salinity-normalized  $A_T$  with SST [Millero et al., 1998] was derived using subsets ( $n = 1,740$ ) of historical  $A_T$  data as well as those collected during the global carbon survey of the 1990s. Since the global  $A_T$  relationships first became available, a significant number of new  $A_T$  measurements have been added to the global data set. Recently, Lee et al. [2006] used surface  $A_T$  measurements ( $n = 5,692$ ) from the Global Ocean Data Analysis Project (GLODAP) v1.1 data set [Key et al., 2004] to determine the relationships of  $A_T$  with SSS and SST for different ocean regimes. A function of SSS and SST in the form

$$A_T = a + b(\text{SSS} - 35) + c(\text{SSS} - 35)^2 + d(\text{SST} - 20) + e(\text{SST} - 20)^2$$

fits  $A_T$  data for each of five oceanographic regimes within an area-weighted uncertainty of  $\pm 8.1 \mu\text{mol kg}^{-1}$  ( $1\sigma$ ).

The global distribution of surface  $A_T$  for August calculated using the derived  $A_T$  algorithms [Lee et al., 2006] applied to climatological SSS and SST fields [Conkright et al., 2002] is shown in

Figure 1. The surface  $A_T$  field shows broad maxima in the central gyres centered at about  $25^\circ$  north and south of the Equator, similar to the pattern for salinity. Poleward of the tropics and subtropics, the annual excess precipitation over evaporation increases, and thus salinity decreases, with latitude. As a result, the values of  $A_T$  generally decrease with increasing latitude. Seasonal changes in the intensity of the convective mixing here are additional key factors that act to change surface  $A_T$  in a measurable way. During seasonal cooling, the intensive vertical mixing brings deep waters rich in  $\text{CaCO}_3$ -driven  $A_T$  to the surface, and thus increases surface  $A_T$ . During seasonal warming, however, the shoaling of the mixed layer makes the contribution of  $A_T$ -rich deep waters to the change in surface  $A_T$  minimal.

The magnitude of seasonal  $A_T$  variability (Figure 2) is directly proportional to the seasonal variations in salinity. Seasonal  $A_T$  variations are generally larger in the tropics and subtropics than in the higher-latitude oceans. In particular, larger amplitudes of seasonal variability are observed in areas where freshwater inputs through rivers (e.g. near the Amazon River and the Bay of Bengal) and the melting of ice (e.g. near sea-ice edges) are significant, or where tropical upwelling occurs.

Finally, since the five regional  $A_T$  algorithms presented in Lee et al. [2006] were derived from the most comprehensive surface  $A_T$  data available, we propose that this new set of  $A_T$  equations should be used for the prediction of surface  $A_T$  over large oceanic areas where SSS and SST values are known.

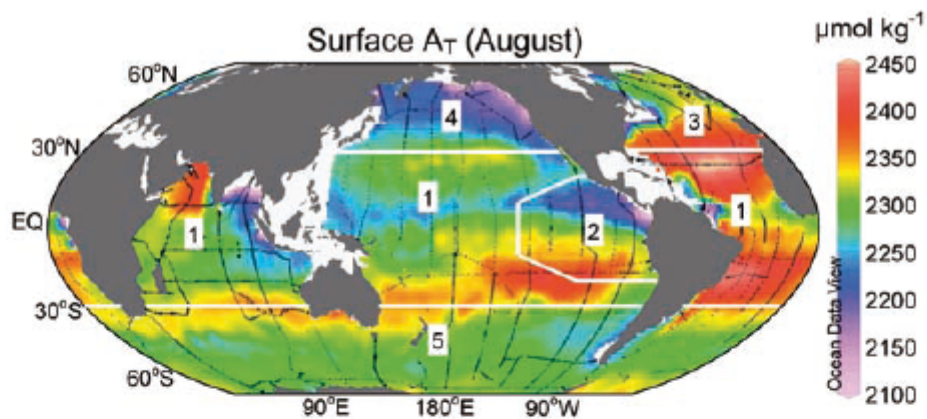


Figure 1. Global distribution of surface  $A_T$  ( $\mu\text{mol kg}^{-1}$ ) for August for the world's ocean. White lines represent the approximate boundaries for the five different regions with unique  $A_T$  algorithms and the numerical numbers represent the five regions. Crosses denote locations of sampling stations.

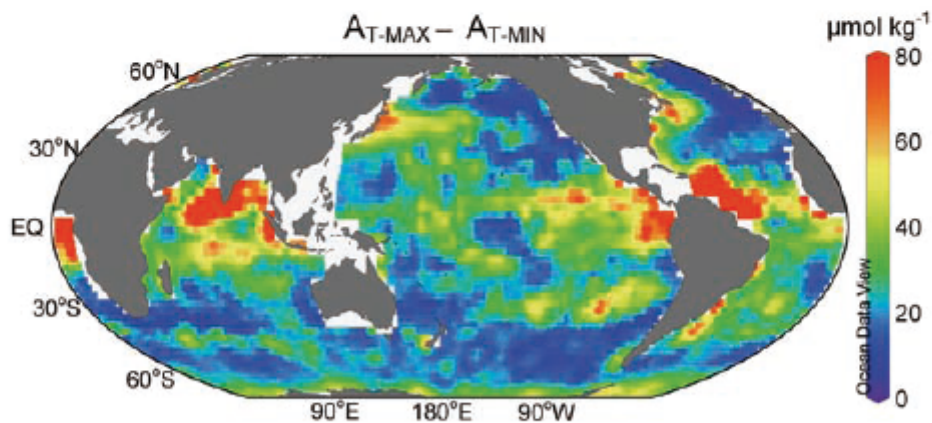


Figure 2. Distribution of the seasonal amplitude (maximum  $A_{TMAX}$  – minimum  $A_{TMIN}$ ) of surface  $A_T$  predicted from the regional  $A_T$  relationships given in Lee et al. [2006], applied to monthly mean sea surface salinity and temperature fields.

#### References

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